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Radiometric dating of the United Kingdom SWAP sites

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Measurements of ²¹⁰Pb by direct gamma assay have been used to date sediment cores from Surface Water Acidification Project (SWAP) study sites in the U.K. The results were checked against additional dating evidence from the artificial fallout isotopes ¹³⁷Cs and ²⁴¹Am. At one of the sites, Devoke Water in Cumbria, the ¹³⁷Cs and ²⁴¹Am data were crucial in identifying a recent sediment hiatus. At sites with recently afforested catchments the sediment record indicated substantial increases in accumulation rates.

Introduction

Accurate dating techniques are crucial to the use of sediment records for palaeolimnological reconstructions of recent environmental change. The precision to which this can be accomplished has been greatly enhanced over the past 20 years by the development of a range of radiometric techniques that use radioisotopes both from natural and artificial sources. The successful utilization of these methods depends on a clear understanding of the mechanisms by which the radioisotopes are incorporated into the sediments. The principal isotope for dating on the timescale of the past 100-150 years, the period of greatest relevance to the Surface Water Acidification Project (SWAP) Palaeolimnology Programme, is 210 Pb. By using this isotope, determination of lake sediment dates with a precision of five to ten years is often attainable. Artificial radioisotopes such as ¹³⁷Cs and ²⁴¹Am have been present in the environment only for the past 35 years and can be used for dating only the most recent sediments. The use of these isotopes does however provide a valuable check on the accuracy of the ²¹⁰Pb calculations. By using these techniques, our principal objective within the SWAP programme was to provide a reliable sediment chronology at each site. We report here on the results of this project.

Methods

Sediments from each core were measured non-destructively for ²¹⁰Pb, ²²⁶Ra, ¹³⁷Cs and ²⁴¹Am by direct gamma assay by using an Ortec HPGe GWL series well-type coaxial low background intrinsic germanium detector (Appleby et al. 1986). Background suppression for this detector is achieved by using a 100 mm thick lead castle, a 305 mm diameter × 305 mm long sodium iodide (NaI(Tl)) escape suppression shield and a 3 mm thick copper lining. ²¹⁰Pb is measured by its gamma emissions at 46.5 keV and 226Ra by the 295 keV and 352 keV γ rays emitted by its daughter isotope, ²¹⁴Pb; ¹³⁷Cs and ²⁴¹Am are measured by their emissions at 662 keV and 59.5 keV, respectively. The absolute efficiency of the detector has been measured by using a series of calibrated sources and sediment samples of known activity. The effect of self absorption of low energy γ rays has been estimated by using sources of different masses. Background counts are done at regular intervals to ensure maintenance of the low background characteristics.

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²¹⁰Pb dating

²¹⁰Pb occurs naturally as one of the radioisotopes in the ²³⁸U decay series. Radioactive disequilibrium between ²¹⁰Pb and its parent isotope, ²²⁶Ra (half-life 1600 years), arises through the mobility of the intermediate gaseous isotope ²²²Rn. A proportion of the ²²²Rn formed by ²²⁶Ra decay in soils diffuses into the atmosphere where it decays to ²¹⁰Pb. This is precipitated onto the land surface and into lakes where it is adsorbed onto sedimentary particles. Pathways by which ²¹⁰Pb accumulates in lake sediments are discussed in detail in Oldfield & Appleby (1984). ²¹⁰Pb activity in sediments in excess of the fraction that derived from decay of the in situ ²²⁶Ra is called unsupported ²¹⁰Pb. It declines in accordance with the exponential radioactivedecay law and can be used for age determinations provided there is an appropriate model for estimating the initial activity. The 210Pb half-life of 22.26 years makes it well suited to dating sediments laid down over the past 100-150 years. Unsupported 210Pb is measured by subtraction of ²¹⁰Pb supported by the parent ²²⁶Ra from the total ²¹⁰Pb activity. In most situations the supported ²¹⁰Pb can be assumed to be in radioactive equilibrium with the in situ ²²⁶Ra. Figure 1 a shows the total and supported ²¹⁰Pb activity against depth for a sediment core from Loch Chon in the Trossachs, Central Scotland. The unsupported ²¹⁰Pb activity in this core is shown in figure 1b.

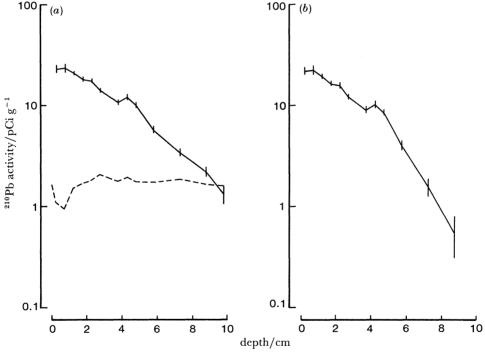


FIGURE 1. ²¹⁰Pb activity against depth in a core from Loch Chon, Scotland; (a) total (——and supported (——) ²¹⁰Pb and (b) unsupported ²¹⁰Pb.

There are two principal models for determining the initial activity of a sediment and hence for calculating ²¹⁰Pb dates, the CRS (constant rate of unsupported ²¹⁰Pb supply) model and the CIC (constant initial ²¹⁰Pb concentration) model. Variants of these models have been developed to account for processes such as sediment mixing. The CRS model (Appleby & Oldfield 1978;

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Robbins 1978) is perhaps the most widely accepted. It is based on the hypothesis that the ²¹⁰Pb supply is dominated by a constant direct atmospheric fallout. There is some evidence (Krishnaswamy & Lal 1978) that the atmospheric ²¹⁰Pb flux may be subject to short-term fluctuations, but as most sediment samples span several years, short-term variations will generally be smoothed out. This model will not be valid where there are alterations or interruptions to the ²¹⁰Pb supply due to, for example, sediment focusing or a sediment hiatus. In these cases dates are calculated either by the cic model (where there is evidence for a constant primary accumulation rate) or by using a composite of both models. Factors governing model choice are set out in Appleby & Oldfield (1983) and Oldfield & Appleby (1984). In this study, the CRS model has generally been used, though dates have routinely been calculated by both models and elements of the cic model incorporated where this has been thought appropriate.

¹³¹Cs and ²⁴¹Am dating

Although ²¹⁰Pb is now routinely used for dating, problems frequently arise over the interpretation of data from sites with disturbed sediment records. In these cases, dates given by the artificial isotope ¹³⁷Cs may be of considerable value. Until the Chernobyl accident in 1986, the most common source of ¹³⁷Cs was fallout from the atmospheric testing of nuclear weapons. Where this isotope is strongly adsorbed onto sediments, the variation of the ¹³⁷Cs activity with depth in a core should reflect the fallout history, with the onset of fallout in 1954 and peak fallout in 1963 providing distinct chronological markers. These may be used to confirm ²¹⁰Pb dates where they are unambiguous, or to resolve ²¹⁰Pb dates where they are uncertain. In practice the value of ¹³⁷Cs dating has often been significantly reduced by the evident mobility of this isotope (Davis *et al.* 1984). The problem has recently been exacerbated by the fallout

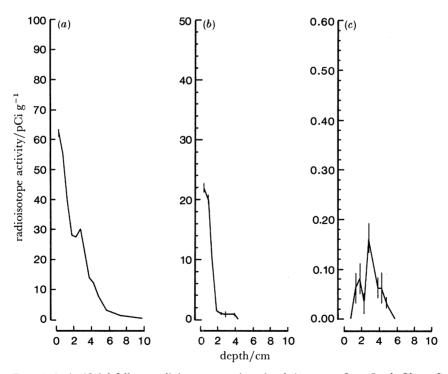


Figure 2. Artificial fallout radioisotopes against depth in a core from Loch Chon, Scotland; (a) 137 Cs, (b) 134 Cs and (c) 241 Am.

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of ¹³⁷Cs from the Chernobyl accident. In regions of high fallout, downward diffusion of Chernobyl ¹³⁷Cs has obliterated the weapons-testing ¹³⁷Cs profile. In these cases, measurement of ²⁴¹Am, another fallout product from atmospheric nuclear weapons testing, may provide a useful alternative. The amount of ²⁴¹Am fallout is very small, only about 0.42% of the ¹³⁷Cs fallout (Eakins & Cambray 1985). Despite this, ²⁴¹Am has been detected in cores from many lakes, and evidence from a growing data set suggests that ²⁴¹Am is significantly less mobile than ¹³⁷Cs. Graphs of ¹³⁷Cs, ¹³⁴Cs and ²⁴¹Am activity against depth for the Loch Chon sediment core are shown in figure 2.

RESULTS

²¹⁰Pb chronologies for Loch Chon calculated by using the crs and crc ²¹⁰Pb dating models are illustrated in figure 3. There is good agreement between the two sets of dates for the most recent sediments, both models indicating a reasonably constant accumulation rate since 1954 of 0.013 g cm⁻² a⁻¹. This is supported by the ¹³⁷Cs and ²⁴¹Am data. The ¹³⁷Cs activity derived from weapons testing fallout (calculated by subtracting Chernobyl ¹³⁷Cs from the total ¹³⁷Cs activity) has a well defined peak at 2.75 cm, as does the ²⁴¹Am activity, and this level is dated to 1966 by the crs model and 1964 by the crc model. Both ²¹⁰Pb models indicate an abrupt transition in the early 1950s from an earlier accumulation rate of 0.0096 g cm⁻² a⁻¹ to the above more recent value. A consequence of this transition (Appleby & Oldfield 1978) is that crc model dates below 4 cm are about 10 years younger than the crs model dates. The increase in accumulation rate is presumably associated with the 1950s afforestation programme in the Loch Chon catchment; its onset is indicated by the reduction in unsupported ²¹⁰Pb activity at a depth of 3.75 cm (dated 1953–1956) shown in figure 1b. In view of this direct evidence in support of the crs model assumptions, the crs model dates were thought to be the more reliable.

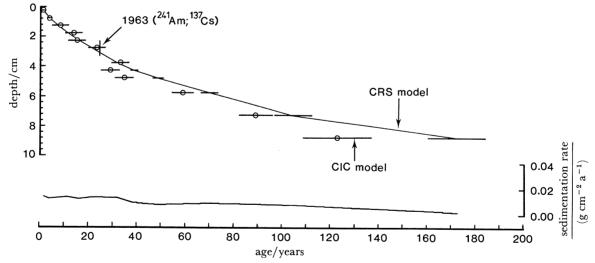


FIGURE 3. Depth against ²¹⁰Pb age and dry mass accumulation rates in a core from Loch Chon, Scotland, as determined by the CRS and CIC ²¹⁰Pb dating models. The graph also shows the depth dated to 1963 by the ¹³⁷Cs and ²⁴¹Am data.

The very high ¹³⁷Cs activity in the top 2 cm of the Loch Chon core can be attributed to fallout from the Chernobyl accident in May 1986. This core was obtained in 1987 and the origin of the ¹³⁷Cs is confirmed by the associated presence of the short-lived isotope ¹³⁴Cs (figure

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2b). Chernobyl fallout was detected in most post-1986 cores and provides evidence of the retrieval of an undisturbed sediment-water interface.

Table 1 gives a summary of the radiometric parameters obtained from SWAP sites in the U.K. Mean accumulation rates are given in table 2. Results from individual sites, summarized below, have been presented in separate reports.

Table 1. Radiometric parameters of SWAP lake-sediment cores

	υ	insupported	²¹⁰ Pb	²²⁶ Ra		
	surf. conc.	invent.	flux	mean conc.	¹³⁷ Cs invent.	²⁴¹ Am invent.
site	(pCi g ⁻¹)	(pCi cm ⁻²)	$(pCi cm^{-2} a^{-1})$	$(pCi g^{-1})$	(pCi cm ⁻²)	(pCi cm ⁻²)
Loch Sionascaig	53.5	12.7	0.40	1.03	10.7^{a}	0.06
Lochan Dubh	10.3	11.3	0.35	1.65	7.9^{a}	0.07
Loch Doilet	6.3	19.3	0.60	1.73	14.4^{a}	0.07
Lochan Uaine	78.1	99.7	3.10	23.06	10.1 ^a	
Coire an Lochan	46.6	16.3	0.51	14.10	3.2^{a}	
Loch Chon	21.9	10.4	0.33	1.66	17.8^{a}	0.04
Loch Tinker	29.6	20.4	0.64	1.33	23.4	0.10
Round Loch	24.9	10.6	0.33	1.47	6.3	0.06
Devoke Water	24.7	34.6	1.68^{b}	0.94	29.3	0.51

^a Denotes sites cored after May 1986.

TABLE 2. COMPARISON OF SEDIMENT ACCUMULATION RATES AT SWAP SITES

	mean accumulation rates								
site	(a) post-1950	(i) $(g cm^{-2} a^{-1})$ (b) 1910–1950		(a) post-1950	$ \begin{array}{ccc} (\text{ii}) & (\text{mm a}^{-1}) \\ (b) & 1910 - 1950 \end{array} $	(b) pre-1910			
Loch Sionascaig	0.0083	0.0086	0.0091	0.75	0.72	0.74			
Lochan Dubh	0.0148	0.0104	0.0074	1.56	0.77	0.47			
Loch Doilet	0.0540	0.0384	0.0186	4.17	2.30	0.97			
Lochan Uaine	0.0223	0.0169	0.0178	0.83	0.52	0.50			
Coire an Lochan	0.0096	0.0096		0.30	0.30				
Loch Chon	0.0133	0.0096	0.0096	1.13	0.52	0.44			
Loch Tinker	0.0193	0.0223	0.0231	1.61	1.47	1.20			
Round Loch	0.0113	0.0120	0.0184	1.51	1.30	1.85			
Devoke Water	0.0525	-		2.90					

Northwest Scotland

²¹⁰Pb dates for Loch Sionascaig were unambiguous and gave a constant sediment accumulation rate over the past 150 years of 0.0088 g cm⁻² a⁻¹ or 0.74 mm a⁻¹. The ¹³⁷Cs results for this core were of no chronological value, even when allowance was made for the contribution from Chernobyl fallout. The ¹³⁷Cs activity was greatest in the topmost sample, and declined monotonically with depth. A small ²⁴¹Am peak was detected at a depth consistent with the ²¹⁰Pb dates.

Loch Doilet and Lochan Dubh both have nonlinear ²¹⁰Pb profiles, suggesting accelerating sediment accumulation rates. The increase has been very much more pronounced in Loch Doilet, presumably because of the disturbances to its catchment arising from the afforestation programme that began in the 1930s. Both cores have relatively well-defined ¹³⁷Cs and ²⁴¹Am peaks at depths reasonably consistent with the ²¹⁰Pb dates. Traces of Chernobyl fallout were observed at the top of the Loch Doilet core.

b 210Pb flux estimated from post-1963 inventory.

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The Trossachs

Despite its much smaller catchment, Loch Tinker has experienced significantly higher sediment accumulation rates than Loch Chon. However, whereas the core from Loch Tinker has a linear ²¹⁰Pb profile suggesting reasonably constant accumulation rates over the past 150 years, the data from Loch Chon indicates accelerated accumulation rates over the past 35 years, presumably due to afforestation in the early 1950s. In both cores the 210 Pb dates are supported by ¹³⁷Cs and ²⁴¹Am data. The enhanced ¹³⁷Cs activities at the top of the Loch Chon core derive from Chernobyl fallout and have virtually obliterated the 1963 peak caused by weapons testing fallout.

The Cairngorms

Lochan Uaine and Coire an Lochan are both characterized by very high unsupported ²¹⁰Pb activities in the topmost sediments. This usually indicates very slow sediment accumulation rates, though at both these sites the very high 226Ra activities may be a contributory factor. In the Lochan Uaine core ²¹⁰Pb equilibrium (corresponding to ca. 150 years of sediment accumulation) is achieved at 8 cm. In Coire an Lochan it appears that equilibrium is achieved at only 2.25 cm. The 210 Pb dates for Lochan Uaine suggest a modest increase in sediment accumulation since 1970. This is supported by ¹³⁷Cs data.

Galloway and Cumbria

The ²¹⁰Pb data from Round Loch of Glenhead indicates constant sedimentation since ca. 1900. There is some suggestion of higher accumulation rates in the 19th century, though this has not been observed in other cores from this lake. The ²¹⁰Pb dates are consistent with the ¹³⁷Cs and ²⁴¹Am data. The data from Devoke Water is contradictory in that ²¹⁰Pb equilibrium appears to have been achieved at 12.5 cm, though the ¹³⁷Cs and ²⁴¹Am data point clearly to an early 1960s date for sediments at a depth of 7 cm. This incompatibility is explained by a sediment hiatus at 9 cm, dated 1954. The high surface concentrations coupled with the high inventories (table 1) suggest that the primary sedimentation rate is quite low but that the core is from a site of intensive sediment focussing. The ²¹⁰Pb and ¹³⁷Cs results suggest that the primary sedimentation rate is reasonably constant. On this assumption the ²¹⁰Pb indicates a loss of ca. 50 years from the sediment record.

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